



#### UNAM-IIMAS

## Joint work with: Panayotis Panayotaros

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A Whitham-Boussinesq long-wave model for variable topography.

Water waves problem with variable depth.





**Discretization of the PDO** associated with the bottom topography.



**Spectral analysis** of the linearized Whitham-Boussinesq model for different families of topographies.



Numerical integration of the evolution of some initial wave-profiles over different topographies.



Work in progress. Project I. II. III. and IV



#### Water waves problem in variable depth

#### Problem setting:



Figure 1. Cartoon of fluid domain

#### EULER'S EQUATIONS

### $\Delta \varphi = 0$ on the fluid domain $N \cdot \nabla \varphi = 0$ on the variable bottom $y = -h_0 + \beta(x)$ Nonlinear boundary conditions on the free surface $y = \eta(x, t)$ Bernullí equation. $\partial_t \varphi + \frac{1}{2} (\nabla \varphi)^2 + g\eta = 0$ Kinematic condition. $\partial_t \eta + \partial_x \eta \cdot \partial_x \varphi - \partial_u \varphi = 0$



## **IDEAL FLUID:** Perfect

- Incompressible
- Irrotational

FLUID DOMAIN: • 2D Simple connected





Hamilton equations with infinitely many  
degrees of freedom:ZAKHAROV  
CRAIG & SULEM  
EQUATIONS
$$\partial_t \begin{pmatrix} \eta \\ \xi \end{pmatrix} = \begin{pmatrix} 0 & I \\ -I & 0 \end{pmatrix} \begin{pmatrix} \frac{\delta H}{\delta \eta} \\ \frac{\delta H}{\delta \xi} \end{pmatrix} \underset{\xi(x,t) = \varphi(x,\eta(x,t),t)}{} \mathcal{E}(x,t) = \varphi(x,\eta(x,t),t)$$
EQUATIONSwith Hamiltonian:  
 $H = \frac{1}{2} \int_{\mathbb{R}} (\xi G(\beta,\eta)\xi + g\eta^2) dx$ 

V.E. Zakharov, 1968, J.W. Miles, 1977, W. Craig and C. Sulem., 1992  $G(\beta,\eta)$  is the DIRICHLET-NEUMANN operator for the fluid domain:

#### **DEFINITION:** Let us consider the solution of th elliptic problem: $\Delta \varphi = 0$ $\varphi(x,\eta(x)) = \xi(x)$

 $N \cdot \nabla \varphi = 0$  On the variable bottom

 $y = -h_0 + \beta(x)$ 

ne
$$[G(eta,\eta)]: \ \xi \longmapsto N\cdot 
abla arphi(| 
abla_x\eta)$$
e bottom







from equibbrium.

This case gives rise to an explicit expression to de DN operator:

 $[G(0,0)]: \xi \longmapsto D \tanh(D)\xi$ 

Where D is as usual the operator  $D = -i\partial_x$ 

However this is not the case for the fluid domain which takes into account the variations in bottom topography as well as the deformations of the free surface



#### Analitic Expansion of operator: $G(\beta, \eta)$

 $G_j(\beta,\eta)$  are homogeneous of degree j in  $\eta$ .

 $G_0(\beta, \eta) = D \tanh(h_0 D) + DL(\beta),$ 

$$G_1(\beta,\eta) = D\eta D$$

$$G_2(\beta,\eta) = \frac{1}{2}(G_0)$$

with  $D = -i\partial_r$ 

#### The bottom variation represented by $\beta(x)$ are taken to be of order O(1), while the surface deformation $\eta(x)$ will be small.

\* Proof for  $G(0,\eta)$  Coifman and Meyer, 1985, \* Craig, Sulem, 2005 \* Craig, Guyenne Nicholls Sulem, 2005 \*Proof for  $G(\beta, \eta)$ , Lannes, For  $||\eta||$ , for  $\beta$  smaller enough (in an appropriate norm)

- $G(\beta,\eta) = G_0(\beta,\eta) + G_1(\beta,\eta) + G_2(\beta,\eta) + \dots$

- $D\eta D G_0 \eta G_0,$   $\frac{1}{2} (G_0 D \eta^2 D D^2 \eta^2 G_0 2G_0 \eta G_1),$

#### and $L(\beta)$ (Involve pseudo-differential operators)



## The operator $L(\beta)$ , can be written in the semi-explicit form: $L(\beta) = -B(\beta)A(\beta),$ where $A(\beta)\xi = \int e^{ikx} \sinh(\beta(x)k) \operatorname{sech}(hk)\hat{\xi}(k)dk,$ $C(\beta)\xi = \int_{\mathbb{T}} e^{ikx} \cosh((-h_0 + \beta(x))k)\hat{\xi}(k)dk,$

and 
$$B(\beta) = C(\beta)^{-1}$$
.

\* W.Craig, P. Guyenne, D. Nicholls & C. Sulem, 2005



#### ADIMENSIONAL PARAMETERS







#### **BOUSSINESQ REGIME**



Aceves-Sánchez, Minzoni and Panayotaros Numerical of a nonlocal Model for water waves with variable depth, 2013



#### **BOUSSINESQ REGIME VARIABLE DEPTH** (SMALL DEPTH VARIATIONS)

 $\epsilon \sim \delta \sim \gamma$ 

#### We want to capture bigger order depth variations!



#### Whitham's type equations

When the full water waves problem is linearized around the zero solution

$$\begin{cases} \eta_t = G(0,\eta)\xi\\ \xi_t = -g\eta. \end{cases}$$

one finds the classical **DISPERSION RELATION:** 

 $\omega^2 = gk \tanh(h_0 k)$ 

Drawback on Boussinesq equations for our purposes: Poor approximation to full dispersion relation for larger wave numbers. Ill posed problem due to negative sign on larger wave numbers.

amplitudes and the existence of solitons with a cusped profile.

Keep term: 
$$G_0(\beta,\eta) =$$

D. Moldabayev, H. Kalisch and D.Dutykh The Whitham Equation as a Model for Surface Water Waves 2014



- Whitham's- type equations offer the possibility of singularity formation at higher





## topography. Part I.

Constant depth.  $\delta \sim \epsilon \ll 1$ 

 $\tilde{D} = \frac{1}{l}D \qquad \tilde{\eta} = \frac{\eta}{a}$ 

Smooth and small depth variations of order O(ε).  $\delta \sim \epsilon \sim \gamma \ll 1$ 

$$G_{\mathcal{A}_1} = \frac{\tilde{D}}{\sqrt{\epsilon}} \tanh(h_0 \sqrt{\epsilon} \tilde{D})$$

Numerical of a nonlocal Model for water waves with variable depth, Wave Motion 2013

Long-wave approximation of  $G(\beta,\eta)$  in presence of non-trivial bottom

 $H = \frac{1}{2} \int_{\mathbb{R}} \left( \xi G(\beta, \eta) \xi + g \eta^2 \right) dx$ 

 $G_{\mathcal{A}_0} = \frac{\tilde{D}}{\sqrt{\epsilon}} \tanh(h_0 \sqrt{\epsilon} \tilde{D}) + \epsilon \tilde{D} \tilde{\eta} \tilde{D}$ 

## $G_0(\beta,\eta) = D \tanh(h_0 D) + DL(\beta)$

$$-h_0\tilde{D}\tilde{\beta}\tilde{D}+\epsilon\tilde{D}\tilde{\eta}\tilde{D}$$

Aceves-Sánchez, Minzoni and Panayotaros





## <sup>2</sup> Long-wave approximation of $G(\beta,\eta)$ in presence of non-trivial bottom topography. Part II.

Depth variations of order O(1)

A Whitham-Boussinesq long-wave model for variable depth, Wave Motion 2016

Satisfies some structural properties of the exact linear DN operator: •  $G_{\mathcal{A}_2}[\xi](x)$  is real if  $\xi$  is real valued •  $G_{\mathcal{A}_2}$  is a symmetric operator • Spectra of this operator has good asymptotic behavior as κ increase we approach the constant depth dispersion relation an the same condition apply to the eigenfunctions. •  $G_{\mathcal{A}_2}$  is a positive operator.



## $G_{\mathcal{A}_2} = Sym(\frac{\tilde{D}}{\sqrt{\epsilon}} \tanh(\sqrt{\epsilon}(-1+\tilde{\beta}(x))\tilde{D})) + \epsilon \tilde{D}\tilde{\eta}\tilde{D}$

Vargas-Magaña, and Panayotaros









We show in the picture that the only visible effects detected with this model are very attenuated.

Aceves-Sánchez, Minzoni y Panayotaros

Numerical of a nonlocal Model for water waves with variable depth, 2003



## 3D using Higher order expansions of the L( $\beta$ ) operator in powers of $\beta(x)$ ,

P. Guyenne, D. Nicholls Numerical simulations of solitary waves on plane slopes 2007 P.Guyenne, D. Nicholls A high-order spectral method for nonlinear wates waves over moving topography 2007

A possible drawback in this formulation is the presence of higher derivatives in  $L_i$  as j increased and the authors also use high frequency truncations of the derivatives.

In 2007 P.Guyenne and P. Nicholls introduced an accurate numerical method for nonlinear surface water waves for variable bathymetry in 2D and



#### A Whitham-Boussinesq model that involves a pseudo differential operator (PDO).

$$G_{\mathcal{A}_2} = Sym(\frac{D}{\sqrt{\epsilon}} \tanh)$$

Pseudo-differential operator of the form:

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## $H = \frac{1}{2} \int_{\mathbb{R}} (\xi G_{\mathcal{A}_2}(\beta, \eta) \xi + g\eta^2) dx$ $h(\sqrt{\epsilon}(-1+\beta(x))D)) + \epsilon D\eta D$

## $a(x,D)\xi(x) = \int_{\mathbb{D}} a(x,k)\hat{\xi}(k)e^{ikx}dk.$

#### where the function a(x,k) is the symbol of the operator.

## Spectral representation in the 2π-periodic framework $\xi = arphi \mid_\eta$



#### PDO in the $2\pi$ -periodic frame

Let a(x, k) be a function periodic in the variable x of period  $2\pi$ :

$$a(x,D)[\xi](x) = rac{1}{2\pi} \sum_{k=-\infty}^{\infty} a(x,k)$$
 with  $a(x,k)$ 





#### The quadratic form in the Hamiltonian in the $2\pi$ -periodic setting:

$$K_{a(x,D)} = \frac{1}{2} \int_{0}^{t}$$



If  $\xi$  is real valued then  $a(x,D)[\xi]$  is real valued then  $K_{a(x,D)}$  is real valued too!

$$K_{Sym(a(x,D))} = \frac{1}{4\pi} \sum_{[k,\lambda]\in\mathbb{Z}^2} \hat{\xi}_k [\hat{a}_{\lambda-k}(k) + \overline{\hat{a}_{k-\lambda}(\lambda)}] \hat{\xi}_{\lambda}^*$$

 $^{.2\pi}$   $\xi a(x,D) \xi dx$ 

$$\sum_{-\infty}^{\infty} \hat{\xi}_k \hat{\xi}_\lambda \hat{a}_{-k-\lambda}(\lambda)$$
$$-\infty \lambda = -\infty$$



#### Discretization of this PDO associated to the bottom topography.

Galerkin truncations of the quadratic form:

Let 
$$\hat{\xi}^{M} = (\hat{\xi}_{1}, \dots, \hat{\xi}_{M})$$

$$\begin{split} K_{a(x,D)}^{M} &= \frac{1}{4\pi} \sum_{[k_{1},k_{2}] \in J_{M}^{2}} \left[ \hat{\xi}_{k_{1}} \hat{\xi}_{k_{2}} \hat{a}_{k_{1}+k_{2}}^{*}(k_{2}) + \hat{\xi}_{k_{1}}^{*} \hat{\xi}_{k_{2}}^{*} \hat{a}_{k_{1}+k_{2}}(k_{2}) \right] \\ &+ \frac{1}{4\pi} \sum_{[k_{1},k_{2}] \in J_{M}^{2}} \left[ \hat{\xi}_{k_{1}} \hat{\xi}_{k_{2}}^{*} \hat{a}_{-k_{1}+k_{2}}(k_{2}) + \hat{\xi}_{k_{1}}^{*} \hat{\xi}_{k_{2}} \hat{a}_{-k_{1}+k_{2}}^{*}(k_{2}) \right] \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} P^{*} & S \\ S^{*} & P \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} P^{*} & S \\ S^{*} & P \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} P^{*} & S \\ S^{*} & P \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} P^{*} & S \\ S^{*} & P \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} P^{*} & S \\ S^{*} & P \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} P^{*} & S \\ S^{*} & P \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} P^{*} & S \\ S^{*} & P \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} P^{*} & S \\ S^{*} & P \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} P^{*} & S \\ S^{*} & P \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} P^{*} & S \\ S^{*} & P \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} P^{*} & S \\ S^{*} & P \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, (\hat{\xi}^{M})^{*})^{T} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix} \end{pmatrix}$$

$$\begin{split} \mathcal{K}_{a(x,D)}^{M} &= \frac{1}{4\pi} \sum_{\substack{[k_{1},k_{2}] \in J_{M}^{2}}} \left[ \hat{\xi}_{k_{1}} \hat{\xi}_{k_{2}} \hat{a}_{k_{1}+k_{2}}^{*}(k_{2}) + \hat{\xi}_{k_{1}}^{*} \hat{\xi}_{k_{2}}^{*} \hat{a}_{k_{1}+k_{2}}(k_{2}) \right] \\ &+ \frac{1}{4\pi} \sum_{\substack{[k_{1},k_{2}] \in J_{M}^{2}}} \left[ \hat{\xi}_{k_{1}} \hat{\xi}_{k_{2}}^{*} \hat{a}_{-k_{1}+k_{2}}(k_{2}) + \hat{\xi}_{k_{1}}^{*} \hat{\xi}_{k_{2}} \hat{a}_{-k_{1}+k_{2}}(k_{2}) \right] \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M})^{T}, ((\hat{\xi}^{M})^{*})^{T}) \begin{pmatrix} p^{*} & S \\ S^{*} & p \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M-1})^{T}, ((\hat{\xi}^{M-1})^{*})^{T}) \begin{pmatrix} p^{*} & S \\ S^{*} & p \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M-1})^{*} \end{pmatrix}, \end{split} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M-1})^{T}, ((\hat{\xi}^{M-1})^{*})^{T}) \begin{pmatrix} p^{*} & S \\ S^{*} & p \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M-1})^{*} \end{pmatrix}, \end{aligned} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M-1})^{T}, ((\hat{\xi}^{M-1})^{*})^{T}) \begin{pmatrix} p^{*} & S \\ S^{*} & p \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M} \\ (\hat{\xi}^{M-1})^{*} \end{pmatrix}, \end{aligned} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M-1})^{T}, ((\hat{\xi}^{M-1})^{*})^{T}) \begin{pmatrix} p^{*} & S \\ S^{*} & p \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M-1} \\ (\hat{\xi}^{M-1})^{*} \end{pmatrix}, \end{aligned} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M-1})^{T}, ((\hat{\xi}^{M-1})^{*})^{T}) \begin{pmatrix} p^{*} & S \\ S^{*} & p \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M-1} \\ (\hat{\xi}^{M-1})^{*} \end{pmatrix} \end{pmatrix}, \end{aligned} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M-1})^{T}, ((\hat{\xi}^{M-1})^{*})^{T}) \begin{pmatrix} p^{*} & S \\ S^{*} & p \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M-1} \\ (\hat{\xi}^{M-1})^{*} \end{pmatrix} \end{pmatrix}, \end{aligned} \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M-1})^{T}, ((\hat{\xi}^{M-1})^{*})^{T}) \begin{pmatrix} p^{*} & S \\ S^{*} & p \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M-1} \\ (\hat{\xi}^{M-1})^{*} \end{pmatrix} \end{pmatrix},$$
 \\ &= \frac{1}{4\pi} ((\hat{\xi}^{M-1})^{T}, ((\hat{\xi}^{M-1})^{\*})^{T}) \begin{pmatrix} p^{\*} & S \\ S^{\*} & p \end{pmatrix} \begin{pmatrix} \hat{\xi}^{M-1} \\ (\hat{\xi}^{M-1})^{\*} \end{pmatrix} \end{pmatrix},







 $K_{a(x,D)}^{M} = \frac{1}{4\pi} (\theta^{T}, \zeta^{T}) \begin{pmatrix} Re(P^{*} + S) \\ -Im(P^{*} + S) \end{pmatrix}$ Letting  $\mathcal{M} = \frac{1}{4\pi} \begin{pmatrix} Re(P^* + S^* + P + S) & -\frac{1}{4\pi} \\ -Im(P^* + S^* - S - P) & -\frac{1}{4\pi} \end{pmatrix}$ 2 and  $\mathcal{M}_{Sym} = \frac{1}{2}(\mathcal{M} + \mathcal{M}^T)$ ,  $K_{Sym(a(x,D))}^{M} = \left(\theta^{T}, \zeta^{T}\right) \mathcal{M}_{Sym} \left(\frac{\theta}{\zeta}\right).$ 

Letting  $\hat{\xi}^M = \theta + i\zeta, \theta, \zeta \in \mathbb{R}^M$ , I Splitting of  $\xi$ -fourier coefficients in real variables

$$S^* + P + S) = -Im(P^* - S^* + S - P) \\ S^* - S - P) = -Re(P^* - S^* - S + P) \begin{pmatrix} \theta \\ \zeta \end{pmatrix}$$

Symmetrization



## Мх2М



#### **5** Spectral analysis of the $M_{sym}$ matrix for different families of topographies.



#### **Bottom Topography**





The matrix corresponding to  $\beta=0$  is a diagonal matrix

 $K_{Max} = 2^9$ ε =.01





















The eigenfrequencies decrease with bathymetry. There is evidence of monotonicity in the Curve with the depth. When you increase the amplitude on the depth, all the eigenvalues goes down. And in the limit as k you go to infinity we have that we approach the constant depth dispersion relation





### **Steepening I.**



The spectral analysis of matrices pointing to show evidences that the depth variation produce significant effects on the eigenmodes:

### **Steepening II.** Topography leads to steeper profiles!





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### Modulation





### Mode 15 $k_{Max} = 2^7$



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#### **Band width**





#### ¿Do there exist a 3-wave resonance?

$$\omega_{k_1} + \omega_{k_2} = k_1 + k_2 = k_2 + k_2 = k_1 + k_$$

Using the Bathymetry can we have solutions of this kind for not small wave number values?





#### Numerical integration of the evolution of some initial wave-profiles over different topographies.

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$$H_2 = \frac{1}{2} \int_{\mathbb{R}} \left[ \xi(Sym(\frac{D}{\sqrt{\epsilon}} \tanh \theta)) \right]_{\mathbb{R}}$$

$$\begin{cases} \partial_t \hat{\eta}_k = \frac{\partial H}{\partial \hat{\xi}_k^*} \\ \partial_t \hat{\xi}_k = -\frac{\partial H}{\partial \hat{\eta}_k^*} \end{cases}, \quad k \end{cases}$$

• 
$$a_k = Re(\eta_k)$$

• 
$$\beta_k = Im(\eta_k)$$
  $d$ 

• 
$$\gamma_k = Re(\hat{\xi}_k)$$
  $dt$ 

• 
$$\delta_k = Im(\xi_k)$$

#### FULLY SPECTRAL EQUATIONS.

 $h(\sqrt{\epsilon}h(x)D)) + \epsilon D\eta D)\xi + \eta^2 dx,$ 

### $k \in J_M$ , with $J_M = [1, ..., M]$ .

$$\begin{pmatrix} \alpha \\ \beta \\ \gamma \\ \delta \end{pmatrix} = \begin{pmatrix} g_1(\alpha, \beta, \gamma, \delta) \\ g_2(\alpha, \beta, \gamma, \delta) \\ g_3(\alpha, \beta, \gamma, \delta) \\ g_4(\alpha, \beta, \gamma, \delta) \end{pmatrix}$$

FOURTH-FIFTH ORDER ADAMS-BASHFOR/MOULTON

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## Second-order approximation of Stokes wavetrain $k_{Max}=2^{5}_{dt=0.001}_{\epsilon=0.01}$



$$\eta_0(x) = a\cos(\lambda x) + \mu_2 a^2 \cos(2\lambda x),$$
  

$$\xi_0(x) = \nu_1 a \cosh(\lambda(\eta_0 + h)) \sin(\lambda x) + \nu_2 a^2 \cosh(2\lambda(\eta_0 + h)) \sin(2\lambda h))$$

$$\mu_2 = \frac{1}{2}\lambda \coth(h\lambda) \left(1 + \frac{3}{2\sinh(\lambda h)}\right),$$
$$\nu_1 = \frac{\omega}{\lambda\sinh(h\lambda)}, \quad \nu_2 = \frac{3}{8}\frac{3\omega}{\sinh^4(h\lambda)}.$$

 $a = 0.065, \lambda = 5.$ 

Craig and Sulem, Numerical Simulation of gravity waves.,1992,





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## the dynamics of surface waves.

Our numerical simulations using the Whitham–Boussinesq model suggest that variable depth has significant effects on



### Second-order approximation of Stokes wavetrain

 $k_{Max}=2^5$ 10.0=3 dt= 0.001  $t_{final} = 120$ 



$$\eta_0(x) = a\cos(\lambda x) + \mu_2 a^2 \cos(2\lambda x),$$
  

$$\xi_0(x) = \nu_1 a\cosh(\lambda(\eta_0 + h))\sin(\lambda x) + \nu_2 a^2 \cosh(2\lambda(\eta_0 + h)))$$

$$\mu_2 = \frac{1}{2}\lambda \coth(h\lambda) \left(1 + \frac{3}{2\sinh(\lambda h)}\right),\,$$





## 





$$\xi(\tau) = \varepsilon \frac{-1}{k^2} \sin(k\tau) + \varepsilon^2 \frac{-B}{4k^2} \sin(2k\tau)$$

c)

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#### Remarks on the accuracy of the numerical integration



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I. Existence of Trapped Modes. Normal modes in a channel of arbitrary cross section. With Panayotaros and Minzoni

II. Comparison of two approaches to the DN operator: PDO approach and CGNS without computing the  $L(\beta)$  operator.

III. The question of whether the particular  $A_{G0}$  or other approximations of the **Dirichlet–Neumann operator that avoid expansions in the depth variation** can be evaluated with an efficiency that is comparable to that of pseudospectral methods.

IV. Looking for triad resonance considering bathymetry.

V Global bifurcation Theorem of the Stokes waves. With Garcia-Azpeitia and Panayotaros.





We are looking for solutions that correspond to an harmonic wave propagating in the x direction without any attenuation or distorsion. The second condition expresses the fact that the transverse energy of the wave is finite and in fact confined, as we shall see later, in some neighbourhood of the coast.

 $\phi($ 



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Figura 1: Stepped Ridge

$$(x, z, y, t) = \operatorname{Re}(\psi(z, y)e^{i(\omega t - \beta z)}) = \psi(z, y)\cos(\omega t - \beta z),$$

$$\int_{\Omega} (|\psi(z,y)|^2 + |\nabla\psi(z,y)|^2) dx dy < +\infty$$





### Constant Depth

Variable Depth

 $\left[\sqrt{\beta_0^2 + \mathbf{D}^2} \tanh(h(X)) \sqrt{\beta_0^2 + \mathbf{D}^2}\right]$ 

# $\begin{cases} -\Delta \psi = -\beta^2 \psi \text{ en } \Omega, \\ \frac{\partial \psi}{\partial y} = \overline{f} \text{ en } \Omega \cap \Gamma_F, \\ \frac{\partial \psi}{\partial n} = 0 \text{ en } \Omega \cap \Gamma_B \end{cases}$

 $\omega^2 \psi(z,0) = f$ 

 $G_{\mathcal{A}_0}(\beta)f = \left[\sqrt{\beta_0 + \mathbf{D}^2} \tanh(h_0\sqrt{\beta_0 + \mathbf{D}^2})\right]f$ 





#### Figura 4: Flatted Shelf with five different highs



Figura 10: Even and Odd Trapped Modes for n=2 and n=3 with  $\beta_0 = 2,15\pi$ 



#### Normal modes in a channel of arbitrary cross section.







A symmetric triangular channel

Figura 11: Symmetric triangular channel

Mark D. Groves HAMILTONIAN LONG-WAVE THEORY FOR WATER WAVES IN A CHANNEL 1994

#### II. Spectral comparison of the *two operators approaches* to the DN operator 1. By Craig, Guyenne Nicholls Sulem and 2. The PDO approach.

We would like to compare the spectrum of:

## 1. The operator derived from the expansion of the DN operator by Craig, Guyenne, Nicholls and Sulem:

$$\mathcal{O}p_{CS} = \frac{1}{\sqrt{\delta}} (D \tanh(\sqrt{\delta}D)\xi - \frac{1}{\sqrt{\delta}} D(B_{\delta}[\gamma\beta]A_{$$

2. The operator involving the PDO approach.

$$\mathcal{A}_{G_0} = \frac{1}{\sqrt{\delta}} D \tanh(\sqrt{\delta}h(x)D) = \frac{1}{\sqrt{\delta}} D \tanh($$

 $\beta])\xi$ 

$$A(\beta)\xi = \int_{\mathbb{R}} e^{ikx} \sinh(\beta(x)k) \operatorname{sech}(hk)\hat{\xi}(k)dk,$$
  

$$B(\beta) = C(\beta)^{-1}.$$
  

$$C(\beta)\xi = \int_{\mathbb{R}} e^{ikx} \cosh((-h_0 + \beta(x))k)\hat{\xi}(k)dk,$$

 $\overline{\delta}(1-\beta(x))D)$ 



#### By solving the generalized eigenvalue problem presented below

$$\frac{1}{\sqrt{\delta}}((D\tanh(\sqrt{\delta}D) - D(B_{\delta}[\gamma\beta]A_{\delta}[\gamma\beta]))\xi =$$

$$((D \tanh(\sqrt{\delta}D) - D(B_{\delta}[\gamma\beta]A_{\delta}[\gamma\beta]))\xi =$$

$$((\tanh(\sqrt{\delta}D) - (B_{\delta}[\gamma\beta]A_{\delta}[\gamma\beta]))\xi =$$

$$((C \tanh(\sqrt{\delta}D) - \beta]A_{\delta}[\gamma\beta]))\xi =$$

$$= \lambda \xi$$
$$= \sqrt{\delta} \lambda \xi$$
$$= \sqrt{\delta} \lambda D^{-1} \xi$$
$$= \sqrt{\delta} \lambda C D^{-1} \xi$$











Generalized Dispersive Curve. PDO operator vs CGNS approach.





Eigenmodes PDO op



#### PDO operator vs CGNS approach.





#### III. Global Bifurcation in Stokes waves for the Whitham-Boussinesq equations with Garcia-Aspeitia, Panayotaros

he use the fact that this operator is O(2) equivariant.

- The existence of travelling waves can be setting in a problem where the global Rabinowitz alternative can be applied. Garcia-Azpeitia using an appropriate operator for the Whitham-Boussinesq equations, proves that this
- operator has a global bifurcation in Stokes waves in an appropriate space,



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