Computing Flexural-Gravity Waves in Three Dimensions

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Acknowledgements

This is joint work with

Jean-Marc Vanden-Broeck at University College London

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- Emilian Părău at UEA
- Paul Milewski at University of Bath

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Compute solutions to Euler's equations as efficiently and as accurately as possible.

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Model for Water Waves

For an inviscid, incompressible fluid with velocity potential $\phi(x, y, z, t)$, the forced Euler's equations are given by

$$\begin{cases} \triangle \phi = 0, & (x, y, z) \in \Omega, \\ \phi_z = 0, & z = -h, \\ \eta_t + \eta_x \phi_x + \eta_y \phi_y = \phi_z, & z = \eta(x, y, t), \\ \phi_t + \frac{1}{2} |\nabla \phi|^2 + \frac{1}{F^2} \eta + P(x, y, t) = -D \frac{\delta H}{\delta \eta}, & z = \eta(x, y, t), \end{cases}$$

where

h: depth

F: Froude number

D: flexural rigidity

 $\begin{array}{l} \eta(x,y,t): \text{ variable surface} \\ P(x,y,t): \text{ external pressure distribution} \\ \frac{\delta H}{\delta \eta}: \text{ condition at the interface.} \\ \Omega = \{-\infty < x < \infty, -\infty < y < \infty, -\infty < z < \eta(x,y,t)\} \\ \end{array}$

Models For a Thin Sheet of Ice



We consider two models

Biharmonic (linear) model

$$H_L = D \frac{1}{2} \int (\bigtriangleup \eta)^2 dA$$

Cosserat (nonlinear) model

$$H_N = D \frac{1}{2} \int (\kappa_1 + \kappa_2)^2 dS$$
 with κ_1, κ_2 principle curvatures

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 Switch into surface variables via the velocity potential at the surface

$$q(x, y, t) = \phi(x, y, z = \eta, t)$$

Go into a moving frame of reference

Combine the dynamic and kinematic boundary conditions

Then the steady-state Bernoulli equation becomes

$$\frac{1}{2}\frac{(1+\eta_x^2)q_y^2 + (1+\eta_y^2)q_x^2 - 2\eta_x\eta_yq_xq_y}{1+\eta_x^2 + \eta_y^2} + \frac{\eta}{F^2} + P + D\frac{\delta H}{\delta\eta} = \frac{1}{2}$$

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Models for Ice

The two different models are considered

► Biharmonic (linear) model

$$\frac{\delta H}{\delta \eta} = \nabla^4 \eta$$

Cosserat (nonlinear) model

$$\begin{split} \frac{\delta H}{\delta \eta} &= \frac{2}{\sqrt{a}} \left[\partial_x \left(\frac{1 + \eta_y^2}{\sqrt{a}} \partial_x H \right) - \partial_x \left(\frac{\eta_x \eta_y}{\sqrt{a}} \partial_y H \right) - \partial_y \left(\frac{\eta_x \eta_y}{\sqrt{a}} \partial_x H \right) + \partial_y \left(\frac{1 + \eta_x^2}{\sqrt{a}} \partial_y H \right) \right] \\ &+ 4H^3 - 4KH \end{split}$$

where

$$\begin{split} & a = 1 + \eta_x^2 + \eta_y^2 \\ & H = \frac{1}{2} a^{3/2} \left[(1 + \eta_y^2) \eta_{xx} - 2\eta_{xy} \eta_x \eta_y + (1 + \eta_x^2) \eta_{yy} \right] \\ & K = \frac{1}{a^2} \left[\eta_{xx} \eta_{yy} - \eta_{xy}^2 \right] \end{split}$$

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Reformulation: Boundary Integral Method

Following the formulation by Forbes (1989), use Green's second identity

$$\int_{V} (\alpha \Delta \beta - \beta \Delta \alpha) dV = \oint_{S(V)} \left(\alpha \frac{\partial \beta}{\partial n} - \beta \frac{\partial \alpha}{\partial n} \right) dS$$

where in three dimensions, β is the fundamental solution given by the Green's function

$$\frac{1}{4\pi} \frac{1}{((x-x^*)^2 + (y-y^*)^2 + (z-z^*)^2)^{1/2}}$$

and $\alpha = \phi - x$, which satisfies Laplace's equation.

System of Equations

The final form of equations to solve for flexural-gravity waves in infinite depth is

$$\frac{1}{2} \frac{(1+\eta_x^2)q_y^2 + (1+\eta_y^2)q_x^2 - 2\eta_x\eta_yq_xq_y}{1+\eta_x^2+\eta_y^2} + \frac{\eta}{F^2} + P + D\frac{\delta H}{\delta \eta} = \frac{1}{2}$$
$$\int_{-\infty}^{\infty} \int_{-\infty}^{\infty} \left[(q-q^* - x + x^*)K_1 + \eta_xK_2 \right] dxdy = 2\pi (q^* - x^*)$$

where

$$\begin{split} & \mathcal{K}_1 = \frac{1}{d^{3/2}} (\eta - \eta^* - (x - x^*)^2 \eta_x - (y - y^*)^2 \eta_y) \\ & \mathcal{K}_2 = \frac{1}{d^{1/2}} \end{split}$$

with

$$d(x, y, x^*, y^*, \eta) = (x - x^*)^2 + (y - y^*)^2 + (\eta - \eta^*)^2$$

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Symmetry

Symmetry in y direction

$$\eta(x,y)=\eta(x,-y)$$

 and

$$q(x,y)=q(x,-y)$$

implies additional terms

$$\frac{1}{2} \frac{(1+\eta_x^2)q_y^2 + (1+\eta_y^2)q_x^2 - 2\eta_x\eta_yq_xq_y}{1+\eta_x^2 + \eta_y^2} + \frac{\eta}{F} - \frac{1}{2} = F(\eta)$$
$$\int_0^\infty \int_{-\infty}^\infty \left[(q-q^* - x + x^*)\tilde{K}_1 + \eta_x\tilde{K}_2 \right] dxdy = 2\pi(q^* - x^*)$$

where

$$\tilde{K}_{1} = \bar{K}_{1}(x, y, \eta, x^{*}, y^{*}, \eta^{*}) + \bar{K}_{1}(x, -y, \eta, x^{*}, y^{*}, \eta^{*})$$
$$\tilde{K}_{2} = \bar{K}_{2}(x, y, \eta, x^{*}, y^{*}, \eta^{*}) + \bar{K}_{2}(x, -y, \eta, x^{*}, y^{*}, \eta^{*})$$

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- ▶ Let x_i and y_j be equally spaced points such that i = 1,..., N and j = 1,..., M.
- ▶ Let the vector of unknowns be $q_{\times(i,j)}$ and $\eta_{\times(i,j)}$ such that

$$u = \left[q_{\times(1,1)}, \cdots, q_{\times(N,1)}, \cdots, q_{\times(N,M)}, \eta_{\times(1,1)}, \cdots, \eta_{\times(N,M)}\right]^{T}$$

- Use finite differences to discretise the derivatives
- Obtain 2NM equations

$$G(u)=0$$

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- Use finite differences to discretise the derivatives
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$$G(u)=0$$

To solve the system

- 1. Set up an initial guess u^0
- 2. Until convergence
 - 2.1 Solve $J(u^n)\delta^n = -G(u^n)$
 - 2.2 Set $u^{n+1} = u^n + \lambda \delta^n$, $0 < \lambda < 1$
 - 2.3 Test for convergence

This method relies on an initial guess u^0 and the Jacobian J.

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2.3 Test for convergence

This method relies on an initial guess u^0 and the Jacobian J.

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Jacobian

The sparsity of the linearised Jacobian for flexural-gravity waves



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We consider two ways of solving the system of equations

- 1. Inexact Newton Method: (direct method) uses an inexact Jacobian (not computed at each step).
- 2. Modified Newton Method: (iterative method) using a preconditioned Krylov method to construct the solution, not keeping the full Jacobian matrix.
 - Preconditioner is constructed as shown in Pethiyagoda et al (2014)
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Newton's method is very sensitive to initial conditions. In order to compute different wave amplitudes, generate a bifurcation diagram

- Guess a small amplitude solution
- Use this guess in Newton's method to compute the true solution.
- Scale the previous solution to get a guess for a larger amplitude solution
- Apply Newton's method to find the true solution.

Can use the Jacobians from previous steps in the bifurcation branch as preconditioners.

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Outline

Motivation

Formulation

Numerics

Solutions

Conclusion and Future Work

Forcing Term

We use the following pressure as a forcing for depression waves



Sample Solutions

Solutions for forced waves underneath an ice sheet



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Sample Solutions

Solutions for forced waves underneath an ice sheet



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Bifurcation Branch

Comparison of the bifurcation branches for flexural-gravity waves with the linear and the nonlinear elasticity models

Note: both models give the same wave amplitude, but different Froude numbers

Flexural-Gravity Wave Profiles

Comparison of the solution profiles for linear elasticity model and the nonlinear elasticity model.

Flexural-Gravity Wave Profiles

Comparison of the solution profiles for linear elasticity model and the nonlinear elasticity model.

Flexural-Gravity Bifurcation Branch

Comparison of the bifurcation branch for linear elasticity model and the nonlinear elasticity model.



Elevation waves are represented as crosses and depression waves as circles.

Outline

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Can compute solutions to both models for flexural-gravity waves

 Both models for produce similar shaped profiles, but at different Froude numbers

- The code is easy to use and easy to modify
- A variety of numerical methods have been tested

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Compute accurate free surface waves without a forcing

- Compare the different models quantitatively
- Do free surface depression or elevation waves bifurcate away from 0?

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Thank you for your attention

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